

A simple, but essential **critierion analysis on methan catastrophe** is intoroduced.

(Global year radiative force increase by CH4 eruption)/(melting heat for methan clathrate=MC)=**870**. (melt amount  $\leq$  1GtC).

It's an **amplifier gain=A** by CH4 eruption. So if **feedback loop loss  $B \geq 1/870$** , then evil positive feedback could begin.

If surplus heat(for irrversible process such as melting)partitioning rate into Arctic  $J_B$  be **0.1**, then 99% of  $J_B$  would be allowable not to hit MC, **but 1% hitting could trigger the feedback. If over 3%, the process could be abrupt.**

Then **note** that the criterion depends on the **feedback time lag**, of which estimation has no confidence. The fatal **manmade bad effects** is also no in the consideration. That is, the reality may be more severe.

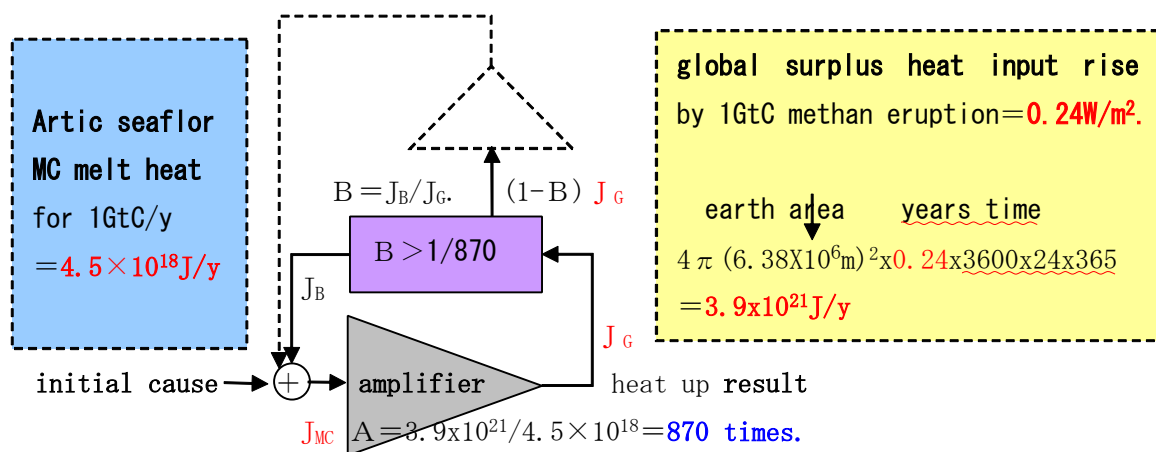
[1]:The General Analysis on Positive Feedback in multi-pathes :

<http://www.geocities.jp/sqkh5981g/FAQ-QL-MC-catastrophe.pdf>

①How much could the **dangerous degree** be estimated for methan catastrophe ?<the primitive circuit model>:

<<from eruption(causing instant golobal heat input rise by the GHG concetration jump) to heat back transfer taward melting target of methane clathrate in Arctic sea flor>>.

①Following paragraphe are **very coarse estimation**, but may be essential. Now let's review on positive feedback process. Result is fed back to enhance cause. For example, methan eruption(by heat input  $J_{MC}$  on MC at sea flor)cause heat input rise on globe  $J_G$ (radiative forcing). Amplified gain  $A = J_G/J_{MC}$ .



②The feedback partitioning(into MC in Arctic sea flor) ratio= $B$ . If  $J_B$  was larger than  $J_{MC}$ , then the system could **run away** without exterior initial cause input.

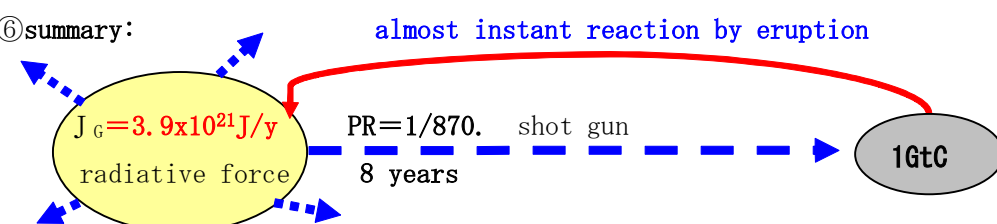
③Note that {heat input into Arctic= $J_B$ /global heat surplus= $J_G$ }  $> 10\%$ .

④Then, if {heat input on MC/heat input into Arctic}  $> 1/87$ , run away could be triggered.

⑤A problem of time delay for feedbacking(heat transfer time into seaflor).

**$0.24 \text{ W/m}^2$**  by 1GtC eruption would be **reserved constant at least 8 years** of methan decay in atmosphere. Therefore time delay could be allowed as within 8 years. Direct solar ray input into arctic sea flor of 200m depth is in a year, while horizontal mean ocean heat transfer from equator to arctic may be a year less than 8 years. Vertical transfer into depth more than 300m may be few in a year.

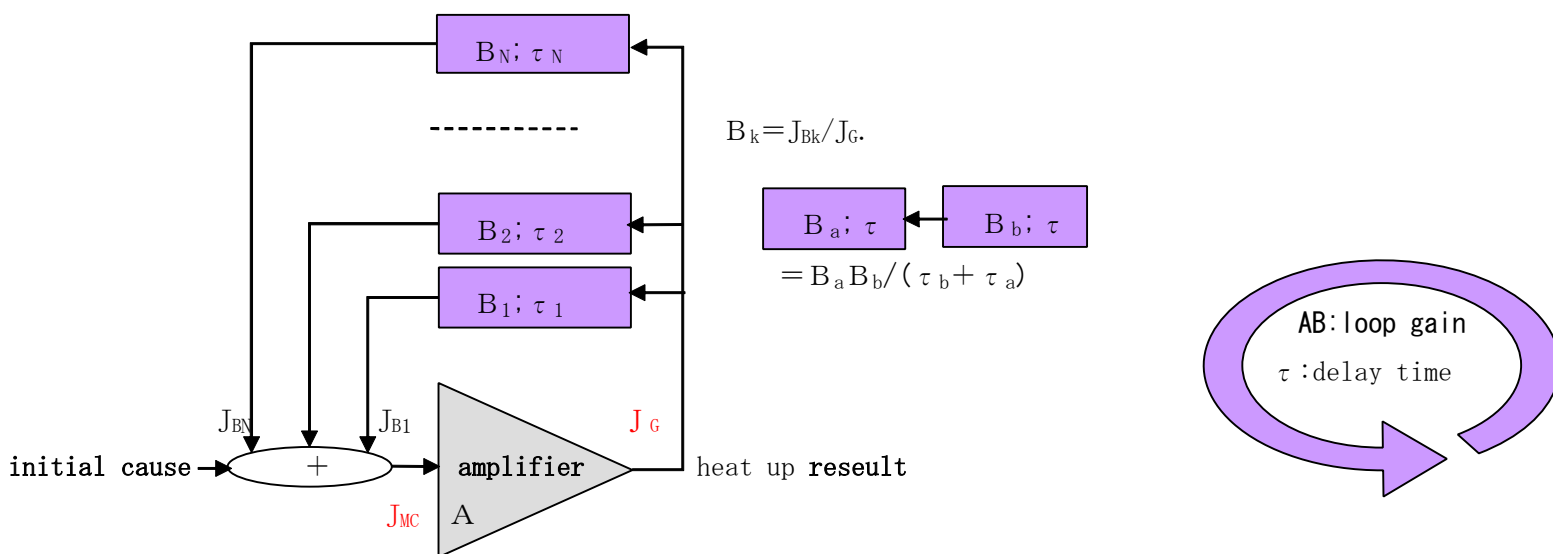
⑥summary:



You could hit out the target with probability 869/870 within 8 years.

②The General Analysis on Positive Feedback in multi-paths :

①circuit model:



②rate equation for of R(t) melting MC amount.

(1)  $dR(t)/dt = R(t) \langle A(t) \sum_{k=1}^N B_k(t) - 1 \rangle / \{ \sum_{k=1}^N B_k(u) \tau_k(t) / \sum_{k=1}^N B_k(u) \}$ .  
 $\longleftrightarrow$  "Time change rate of R=R[loop gain-1]/(mean delay time in looping)"

(2)  $R(t) = R(0) \exp[ \int_0^t du. \langle A(u) \sum_{k=1}^N B_k(u) - 1 \rangle / \{ \sum_{k=1}^N B_k(u) \tau_k(t) / \sum_{k=1}^N B_k(u) \} ]$ .

\*the criterion on becoming positive feedback is  $\{ A(u) \sum_{k=1}^N B_k(u) \geq 1 \}$ .

\*the amplifier gain A is not linear, but is almost linear=870 times below input level 1GtC melt.

(1) **Amplified gain A**  $\equiv J_G / J_{MC}$  is non-linear for input  $J_{MC} = A$  (radiative forcing function of CH4 density), which could be derived following formula in (2) with (5).

(2) **radiative forcing calculator formula for CH3 with NO.**

radiative forcing  $\Delta F$  mean **global heat input rise** ( $W/m^2$ ) by concerned GHG concentration rise of  $M_0$  (1750)  $\rightarrow M$  change. That is, methane eruption  $\Delta M = (M - M_0)$  causes global surplus heat of  $J_G \equiv \Delta Q = 4 \pi R_E^2 \times \Delta F \times 3600 \times 24 \times 365$  in a year (AMP output power). Then note, AMP input power is melting heat for  $\Delta M \times 440 \text{ KJ/Kg} = J_{MC}$ .

<http://ja.wikipedia.org/wiki/%E6%94%BE%E5%B0%84%E5%BC%B7%E5%88%B6%E5%8A%9B>

M=CH4 concentration(ppb), N=that of NO ( $N_0=310$  ppb),

**Methan radiative forcing rise function** by  $M_0 \rightarrow M$  change with  $NO = N_0$ .  
 $\Delta F = 0.036 [\sqrt{M} - \sqrt{M_0}] - \{ f(M, N_0) - f(M_0, N_0) \} \equiv 0.036 [\sqrt{M} - \sqrt{M_0}] - \Delta f$ .

$f(M, N) = 0.47 \ln \{ 1 + 2.01 \times 10^{-5} \times (MN)^{0.75} + 5.31 \times 10^{-15} M (MN)^{1.52} \}$

(3) Methane weight **10GtC** and the concentration=4720ppb,

(4) the initial methane concentration  $M(1750) = 700 \text{ ppb} \equiv M_0$  in 1750.  $M(2000) = 1770 \text{ ppb} = 5 \text{ Gt} (\text{CH}_4=16\text{g}) = 3.75 \text{ Gt} (\text{C}=12\text{g})$  in 2000.

(5) **Methan clathrate melting heat/kg**=440KJ/Kg.  $J_{MC}(10\text{GtC}) = 4.5 \times 10^{18} \text{ J/y}$ .

③Feedback Pathes:

heat path "k"		$B_k = J_{Bk} / J_G$	$\tau_k$
1:	Insolation input on Arctic ocean ( $L_A$ ) $\rightarrow$ MC in Arctic seafloor		instant
2:	Insolatin input into ocnas $\rightarrow$ exterior ocean heat input Aritic (L) $\rightarrow$ MC		$L/v \doteq 0.5 \text{ y}$
3:	Atmospheric heat input into Arctic ocean? $\rightarrow$ MC in Arctic seafloor		
4:	heat output in phase transit of supercooling water? $\rightarrow$ MC in Arctic seafloor	X X X X X X X	

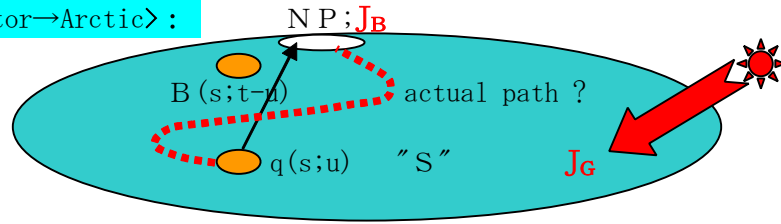
(1) Heat transfer coarse modeling in ocean circulation <  $B_s$  (equator → Arctic) > :

$$J_B = \iint ds \int_0^{t=1} du B_s(s; t-u) q(s; u) = B_s(s*; t-u*) \iint ds \int_0^{t=1} du q(s; u)$$

$$J_B \equiv B_s * (\tau) J_G : \text{feedback energy flow}$$

$$J_G \equiv \iint ds \int_0^{t=1} du q(s; u) : \text{Total insolation input into oceans.}$$

$$B_s(s*; t-u*) \equiv B_s * (\tau).$$



$q(s; u)$  is heat input density on area and time elements  $ds, dt$  from  $J_G$ .  $B_s(s; t-u)$  is heat partitioning rate from position and time  $= (s, u)$  into Arctic at time  $t$ . It's a model of **linear transfer function** (of partitioning rate  $B_s$  with time delay  $\tau$ ).

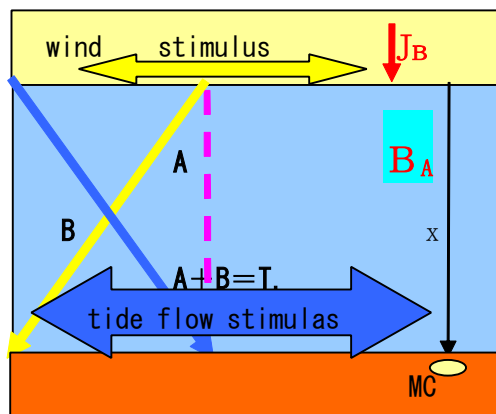
(2) A coarse modeling of heat transfer from ocean upper mixed layer into target MC at seafloor <  $B_A$  (upper NP sea → sea floor) > :

<http://www.geocities.jp/sqkh5981g/PSEUDO-DIFFUSION-BY-TURBULANCE.pdf>

(a) Our concern depth of **Arctic ocean** is limited within 200m~1300m where MC could be. Its rather shallow depth where author assume monotonous distribution of **ocean turbulence** by both tidal flow in deeper and sea surface wind in shallower. An ocean convection with warmer upper layer and colder lower layer never transfer heat from up into bottom.

(b) Pseudo diffusion processes by exponential function distribution  $q(x; t) \equiv (1/\nu_0 t) \exp(-x/\nu_0 t)$ .

→  $\langle x \rangle = 1/\lambda = (\nu_0 t)$ . Average depth at time  $t$  with **37% heat flow passing**.  $\langle x \rangle = 2/\lambda = (2\nu_0 t)$  with **14% hfp**.



(c) The problem of hitting probability on target MC by heat input.

In our scheme of  $B_A$ , it had been already counted in.

More details could be seen in the following.

<http://www.geocities.jp/sqkh5981g/target-hitting-probability.pdf>

(3) Actual observed data:

Those are most weak point in this report. After all, author could not find reliable data for those parameter  $\{B * (\tau)\}$ .

[2]: Conclusion by data **assumed** values.

	A (radiative force gain) = $E_r(M)/E_m(M)$ .	? $B = B_s$ (equator → NP (Arctic)) × $B_A$ (upper NP sea → sea floor).	? $\tau = \tau_s + \tau_A$ .	$(A B - 1) / \tau \equiv K$ .	$\exp[\int_0^1 du. K]$
10MtC	$870 = 3.9 \times 10^{19} \text{J/y} / 4.5 \times 10^{16} \text{J/y}$	$0.10 \times 0.01 = 0.001$	$0.5y + 0.5y$	$(0.87 - 1) / 1 = 0$ .	-
		$0.10 \times 0.012 = 0.0012$	$0.5y + 0.5y$	$(1.0 - 1) / 1 = 0$ .	1
		$0.10 \times 0.03 = 0.003$	$0.5y + 0.5y$	$(2.6 - 1) / 1 = 1.6$	5
		$0.10 \times 0.03 = 0.003$	$0.5y + 1.0y$	$(2.6 - 1) / 1.5 = 1.0$	2.7
		$0.10 \times 0.03 = 0.003$	$0.5y + 1.5y$	$(2.6 - 1) / 2 = 0.8$	2.2
100MtC	$870 = 3.9 \times 10^{20} \text{J/y} / 4.5 \times 10^{17} \text{J/y}$	$0.10 \times 0.01 = 0.001$	$0.5y + 0.05y$	$(0.87 - 1) / 0.5 = 0$ .	-
		$0.10 \times 0.03 = 0.01$	$0.5y + 1.0y$	$(2.6 - 1) / 1.5 = 1.07$	3
1GtC	$870 = 3.9 \times 10^{21} \text{J/y} / 4.5 \times 10^{18} \text{J/y}$	$0.10 \times 0.01 = 0.001$	$0.5y + 0.05y$	$(0.87 - 1) / 0.5 = 0$ .	-
		$0.10 \times 0.03 = 0.03$	$0.5y + 0.5y$	$(2.6 - 1) / 1 = 1.6$	5
10GtC	$580 = 2.6 \times 10^{22} \text{J/y} / 4.5 \times 10^{19} \text{J/y}$				

☞ 全球余剰不可逆化熱の北極海分配が 0.1=10%を仮定、北極入力不可逆化熱 =  $B_A$  の 99%は海洋水温上昇等に消費しても OK、だが 1%が MC に当たると正帰還。勿論海底の何処にどの程度分布かが大問題、なほ北極にも小型ハリケンがあると言う。危険な海洋攪乱に作用、臨界間じかならば Berign strait を閉じてタイムラグを大きくする方法もある。