

—Evolution equation of Global surface Temperature(EGT)with RF—' 9/9/13, 10/20,  
erruntum:Global heat capacity  $C_G$  and the relations are corrected.

Our aim is estimating global temperature trend by GHG concentration change.  
Those factors are (1)constant solar ray input,(2)global radiative forcing(=RF) of atmosphere(GHGs rule trapp(with downward reradiation) and passing cooling radiation from the surface).The radiative ability is function of GHG concentration and albedo.(3)global dynamic surface temperatue(of equivalent global heat capacity).The last factor determine (RF)heat sink amount from (2).

RADIATION FORCING(the general formultion and the intepretation).

<http://www.geocities.jp/sqkh5981g/RADIATION-FORCE.pdf>

[ 0 ] : Global mean temperature as a function of albedo and GHG concentration:

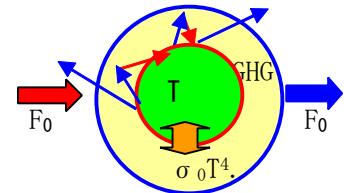
In this chapter,our discussion is asuumed a **thermal equilibrium state**,while discussion on **radiative forcing(RF)** is not such one, but **in-equilibrium**.

①**energy balance equation:input solar heat and cooling radiation output on earth:**

<<(1)Stefan-Boltzmann low>>

$\pi R^2 F_0 (1-m) / (1-b) = 4 \pi R^2 \sigma_0 T^4 \dots (1)$  **energy balance equaiton.**

<<input power=output power>>



$F_0 = 1366 \text{W/m}^2$ =original solar heat input at strat sphere.

$R_E = 6.38 \times 10^6 \text{m}$ =Earth radius, ;  $\sigma_0 = 5.67 \times 10^{-8} \text{Wm}^{-2}\text{K}^{-4}$ .<Stefan Boltzmann>

$m$ =albedo( $\doteq 0.4$ )=reflection(by white clouds, ice sheets)probability for  $F_0$ .

$b$ ( $\doteq 0.48$ )= $1-a$ 、 $a$ =passing probability of cooling radiation through GHG.

" $b$ " is so to say a **super black body factor** rising temperature.

$T = (273 + 15)^\circ\text{C}$ =global mean temperature on **the(almost ocean) surface**.

(2)It's **energy input & output budget equation** at macro global surface of mean

temperature  $T$  with substantial **solar heat input** $=\pi R^2 F_0 (1-m) / (1-b)$ .  $\pi R^2 F_0 (1-m)$  is original solar input onto the surface with **reflection probability m(albedo)**.

The factor  $a=(1-b)$  is passing(heat dissipating) probability into cosmic space of **cooling radiation** from global surface through GHG. $b=(1-a)$  is **trapp and downward re-radiation** probability by GHG. **Boundary**(atmosphere vs global surface) radiative factor **b** is a function of each GHG component concentration.

(3)The output heat of cooling radiation(**infrared ray**) flow rate is  $4 \pi R^2 \sigma_0 T^4$  representing **Stefan-Boltzmann law of radiation field with thermal matters**. It's the equation **in thermal steady equilibrium state**.

## ② Outflow reflection by GHG trapp with and without irreversible RF:

(1) Assumption on complete cooling radiation without radiative forcing.

GHG almostly do'nt trapp input **radiant ray**, but do outflow of **infrared ray**.

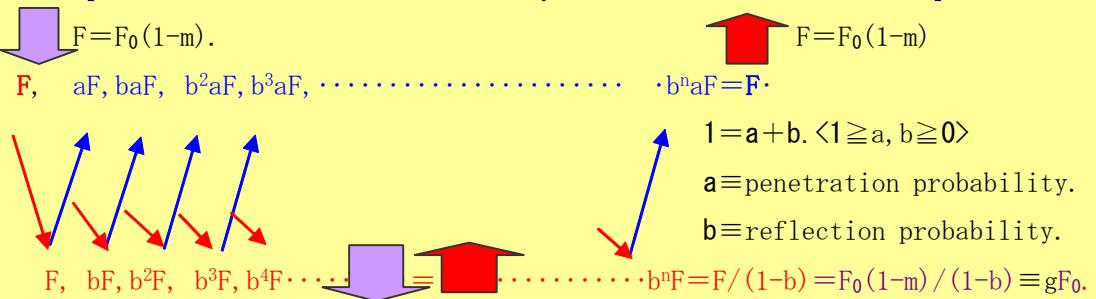
GHG gas on globe enforces infrared ray many times trial to penetrate through

GHG in repeated reflection between globe surface and GHG. We define **ray**

**penetrating probability**  $\equiv a \equiv 1-b$ , then  $F_b$  is energy flow beign trapped and reflecting into the surface by GHG, thus the infinitive times repeating forms so called power series. Then note  $\text{input}=\text{output}=F(=F_0(1-m))$  at top(equibrium).

### GHG EFFECT of EQI1IBRIUM STATISTICAL MODEL

as repeated reflection of infrared ray between surface and atmosphere



<[http://www.fnorio.com/0040Greenhouse\\_effect1/Greenhouse\\_effect1.htm](http://www.fnorio.com/0040Greenhouse_effect1/Greenhouse_effect1.htm)>

$$S_n = F + bF + b^2F + b^3F + b^4F + \dots + b^n F = F/(1-b)$$

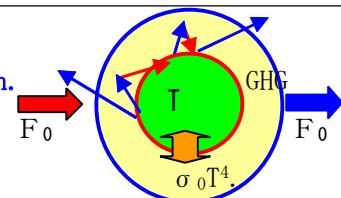
$$R_n = aF + baF + b^2aF + b^3aF + \dots + b^n aF = aF(1 + b + b^2 + b^3 + b^4 + \dots + b^n) = aF/(1-b) = F$$

Thus input flow equivalently increase by GHG**Green House Effect**, or **radiation force** as  $F\{b+b^2+b^3+\dots+b^n+\dots\}$ , where  $b$  is trapping probability. In anyway, major input energy never fail to escape from globe by many many trials.

<<generalized Stefan-Boltzmann law>>

$$\pi R^2 F_0(1-m)/(1-b) = 4 \pi R^2 \sigma_0 T^4 \dots \text{(1) energy balance equaiton.}$$

<<input power=output power>>



Case for  $m=0.4$ .

$T=15^\circ\text{C.} \rightarrow b=0.48, a=0.52$ . <now state>,  $g=1.15$ .

$T=17^\circ\text{C.} \rightarrow b=0.49, a=0.51$ . <critical state>,  $g=1.18$ .

$T(=273+15)^\circ\text{C}=\text{global mean temperature on the(almost ocean) surface.}$

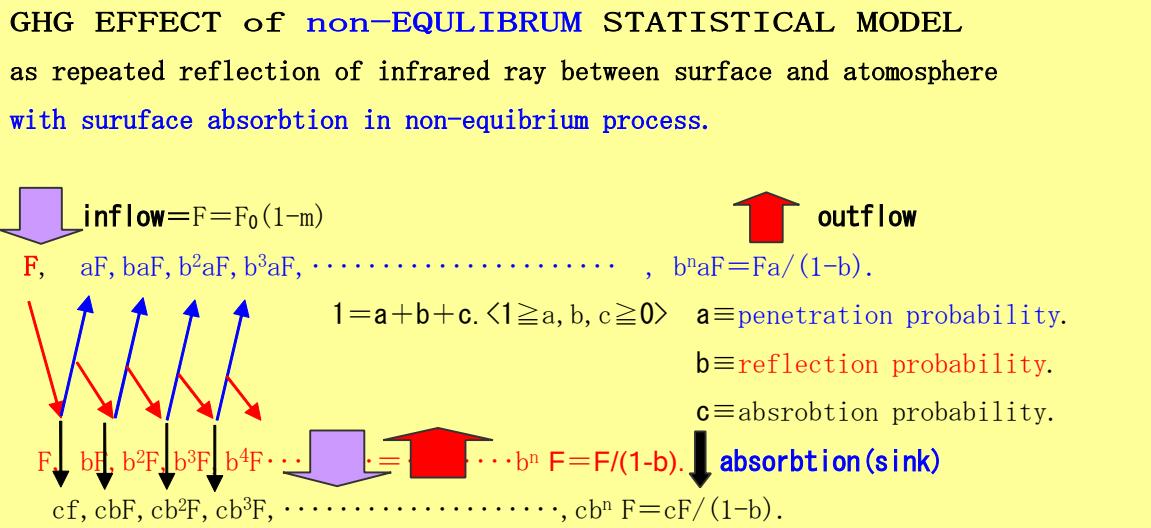
(2) Note that **global mean temperature** is dominated only by two factors= $\{m, b\}$ .

That is **ALBEDO REFLECTION** probability and **GHG TRAPP** probability. This fact is to become main factors also for policy making against climate change crisis.

### ③Incomplete cooling radiation with radiative forcing in "T" transit.

(1) Considering absorbtion(sink) in cooling ray radiation scattering in GHG.

$1=a+b+c \equiv \{\text{penetration}+\text{reflection}+\text{absorbtion}\}$  probability.



(2) The energy conservation for input and output with sink.

Then energy conservation low is input  $F = \{Fa/(1-b) + Fc/(1-b)\} = \text{output} + \text{sink}.$

$$F = \pi R^2 F_0 (1-m) = \pi R^2 F_0 (1-m) a/(1-b) + \pi R^2 F_0 (1-m) c/(1-b).$$

$$\text{Input} = \text{output} + \text{absorbtion}.$$

$$4\pi R^2 \delta F_A(t) \equiv \pi R^2 F_0 (1-m) c/(1-b).$$

### (3) Dynamic global radiative forcing as a heat for absorberes :

$$\delta F_A \equiv \frac{1}{4} F_0 (1-m) c/(1-b). : \text{Dynamic radiative forcing as global absorbtion energy}.$$

Generally to tell, probability "c" is so neglegible(?) small against {a, b}.

However, it's fatally effective on inhabitants on globe.

$$\pi R^2 F_0 (1-m) = \pi (6.38 \times 10^6 \text{m})^2 \times 1366 \text{W/m}^2 \times (1-0.4) = 6.99 \times 10^{16} \text{W}. (\text{bare heat input on globe})$$

$$4\pi R^2 \Delta F_G = 4\pi (6.38 \times 10^6 \text{m})^2 \times 1.6 \text{W/m}^2 = 8.18 \times 10^{14} \text{W}. (\text{IPCC global radiation forcing}).$$

$$4\pi R^2 \sigma_0 T^4 = 4\pi (6.38 \times 10^6 \text{m})^2 \times 5.67 \times 10^{-8} \text{Wm}^{-2} \text{K}^{-4} \times (273+15)^4 = 2.0 \times 10^{17} \text{W}.$$

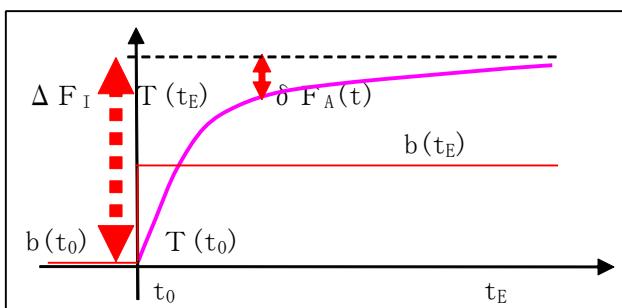
$$K1 \equiv [4\pi R^2 \Delta F_G] / [\pi R^2 F_0 (1-m)] = 0.0117.$$

$$K2 \equiv [4\pi R^2 \Delta F_G] / [4\pi R^2 \sigma_0 T^4] = 0.0041.$$

④ Instantaneous radiative forcing  $\equiv \Delta F_I(t_0; t_E)$  and Actual one  $\equiv \delta F_A(t)$  :

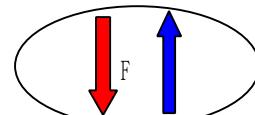
(1) Instantaneous radiative forcing  $\equiv \Delta F_I(t_0; t_E)$ :

Assuming perturbation  $\{b(t_0) \rightarrow b(t_E)\}$  by step function rise of GHG concentration,



we could know the initial and final temperature by ①(1). Then surface equal upward and downward flux could be calculated as follows.  
As temperature rise, the heat sink

would be decreased to zero at last ( $t=t_E$  in equilibrium state).



$$(a) \frac{1}{4}F_0(1-m)/(1-b(t_E)) = \sigma_0 T^4(b(t_E)) \equiv F(t_E). \quad <\text{final equilibrium state}>$$

$$(b) \frac{1}{4}F_0(1-m)/(1-b(t_0)) = \sigma_0 T^4(b(t_0)) \equiv F(t_0). \quad <\text{initial equilibrium state}>$$

$$(c) \Delta F_I(b(t_E), b(t_0)) \equiv F(t_E) - F(t_0) = \frac{1}{4}F_0(1-m) \{1/(1-b(t_E)) - 1/(1-b(t_0))\}.$$

Thus the difference of radiative forcing is a function of  $\{b(t_0), b(t_E)\}$  or GHG concentration variation during  $(t_0, t_E)$ . <<instantaneous RF>>

(2) Dynamic radiative forcing  $\equiv \delta F_A(t)$  the actual one.

$$(a) I = a(t_0) + b(t_0) = a(t_E) + b(t_E) = a(t') + b(t') + c(t'). \quad (t_0 < t' < t_E)$$

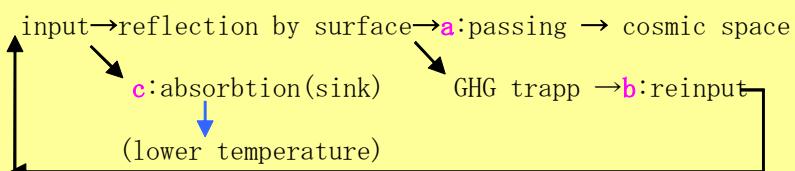
(equilibrium state)                  (non-equilibrium state)

(b) GHG concentration rise shall be increasing  $b(t')$  with  $a(t')$  decreasing and with as  $c(t') > 0$ .

$$I = a(t_0) + b(t_0) = a(t) + b(t) = a(t') + b(t') + c(t'). \rightarrow b(t') + c(t') = 1 - a(t).$$

-(f) radiative forcing as irreversible process in change  $\{b(t_0) \rightarrow b(t_E)\}$  -.

As for an input, the outcomes are {passing, reinput, absorbtion}.



(3) Decision about unknown "c" by known trapp probability "b" or  
change in heat trapp probability "b" and absorbing probability "c":

$$4\pi R^2 \int_{t_0}^{t_E} du. \delta F_A(u) = 4\pi R^2 \int_{t_0}^{t_E} du. c(u) F/(1-b(t_E)) = C_G [T(t_E) - T(t_0)] \equiv \Delta Q_G.$$

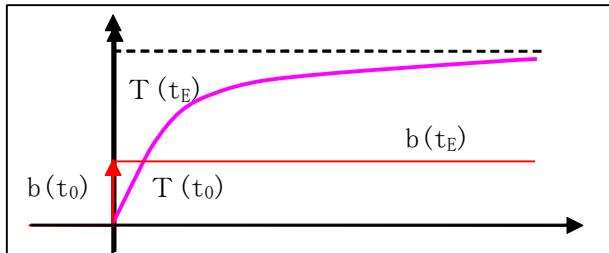
It's a global temperature rise of heat capacity  $C_G$  during  $\{t_0, t_E\}$  due to "c".

(a) absorption probability of time function  $c(t)$  as functional of  $b(t_E)$ .

$$\int_{t_0}^{t_E} du. c(u) = \Delta Q_G(b(t_E), b(t_0)) (1-b(t_E)) / \pi R^2 F_0(1-m).$$

$c(u)$  and also related radiative forcing could not a constant, hence we assume an exponential decay of relaxation process  $\{T(t_0) \rightarrow T(t_E)\}$ .

Assuming perturbation  $\{b(t_0) \rightarrow b(t_E)\}$  by step function rise of GHG concentration,



and now having known initial and final temperature  $\{T(t_0) \rightarrow T(t_E)\}$ , necessary parameter is the global heat capaticity time constant  $\tau_G$ .

$$\rightarrow c(t) \equiv c(t_0) \exp(-(t-t_0)/\tau_G). \rightarrow \int_{t_0}^{t_E} du. c(u) = (c(t_0)/\tau_G) [1 - \exp(-(t_E-t_0)/\tau_G)]$$

$$(c(t_0)/\tau_G) [1 - \exp(-(t_E-t_0)/\tau_G)] = \Delta Q_G(b(t_E), b(t_0)) (1-b(t_E)) / \pi R^2 F_0(1-m).$$

$$* \tau_G = c(t_0) \pi R^2 F_0(1-m) / \Delta Q_G(b(t_E), b(t_0)) (1-b(t_E))$$

$$= c(t_0) \pi R^2 F_0(1-m) / (1-b_E) C_G [F_0(1-m)/4 \sigma_0]^{1/4} [(1-b_E)^{-1/4} - (1-b_0)^{-1/4}].$$

$$F_0(1-m)/4 \sigma_0(1-b) = T^4. ; [T(t_E) - T(t_0)] \equiv [F_0(1-m)/4 \sigma_0]^{1/4} [(1-b)^{-1/4} - (1-b_0)^{-1/4}].$$

$$\Delta Q_G(b(t_E), b(t_0)) = C_G [T(t_E) - T(t_0)] = C_G [F_0(1-m)/4 \sigma_0]^{1/4} [(1-b_E)^{-1/4} - (1-b_0)^{-1/4}].$$

$$\delta F_A(t_0) = c(t_0) F_0(1-m) / (1-b(t_E)) = \frac{1}{4} F_0(1-m) \{1/(1-b(t_E)) - 1/(1-b(t_0))\}.$$

$$* c(t_0) = \frac{1}{4} \{1 - (1-b_E)/(1-b_0)\}.$$

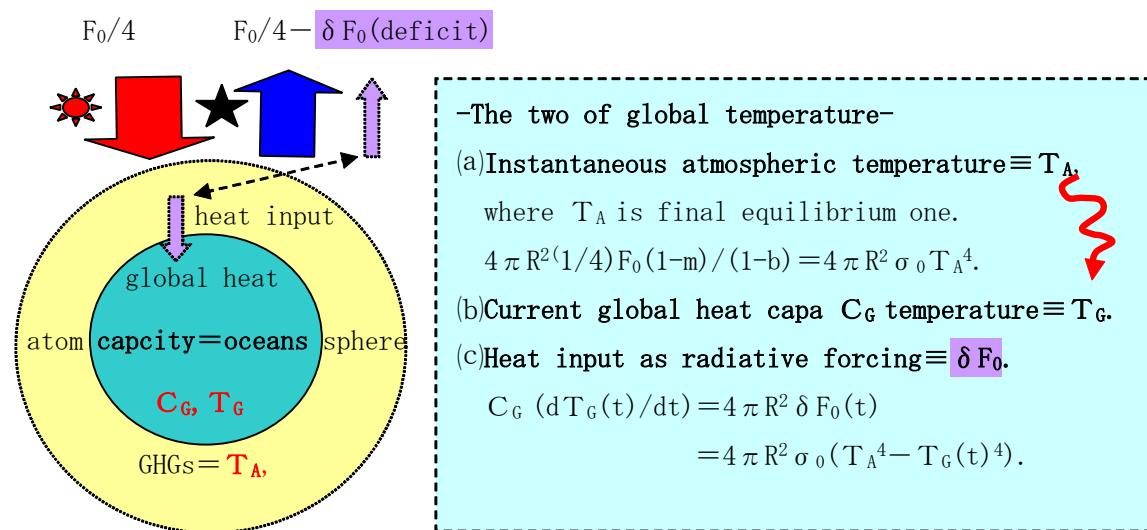
In order to utilize radiative forcing for estimation on temperature rise process, the intensity  $c(t_0) \langle \Delta F_I(t_0; t_E) \rangle$  with time constant  $\tau_G$  are necessary.

As was known in above, discussion at here has not a dynamical principle.

[1] : The heat flow imbalance of global surfaces in macroscopic meaning :

(1)the fundamental scheme:

Our intension is not to mention local weather forcasting, but prediction on long years global climate trend. So it's an averaged something without the abnormaly in time and volume. Consequently macroscopic variable could become available.



(2)Radiative imbalance between global heat capacity and GHG the radiative heat.

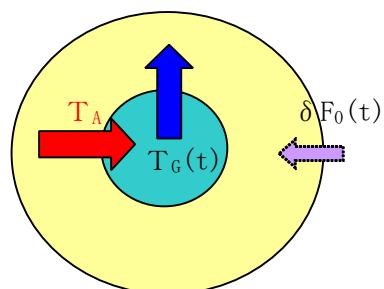
Heat capacity of land and atmosphere is less as 1/1000 of oceans. So it could not be a representative of global temperature. After all, heat deficit at the cosomic gate(strtsphere)  $\delta F_0$  is flow into ocean of heat capacity  $\equiv C_G$ . If  $T_A = T_G$ , dynamic radiative forcing(RF)  $\delta F_0 = 0$  (equilibrium) by balance of radiation of  $T_A$  and  $T_G$ . If  $T_A > T_G$ , dynamic radiative forcing(RF)  $\delta F_0 > 0$  (inequilibrium) due to imbalance. Those **pseudo blackbody radiation imbalance** may be as follows.

$$\delta F_0(t) = \sigma (T_A^4 - T_G(t)^4) \dots \dots \dots (2)$$

(3)Evolution equation of Global surface Temperature(EGT):

$$C_G (dT_G(t)/dt) = 4 \pi R_E^2 \sigma (T_A(t)^4 - T_G(t)^4).$$

$Q_G(t) = C_G T_G(t)$  is global heat amount in global heat capacity  $C_G$ . So its time derivative is heat input per unit time. It must be equal with (inflow-outflow) at the global surface of  $C_G$ . Each flow amount is dominated as **super blackbody radiation** of  $\sigma_0 T^4$  with  $T$ .



#### (4)Temperature of pseudo blackbody radiation for atmosphere:

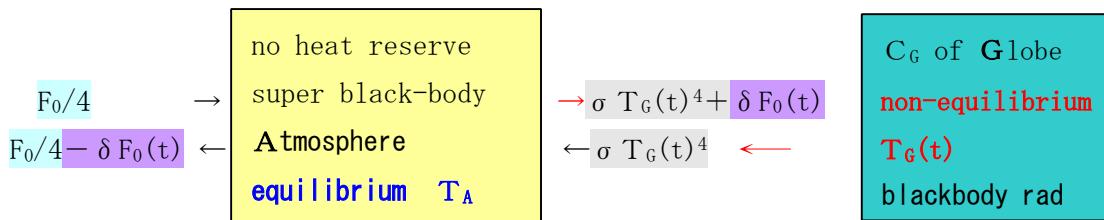
Let imagine radiation flux density of  $F = \sigma T^4$ , of which temperature would be observed as  $T$ .

#### (5)Reexamination: radiative temperature of atmosphere the curiousness :

A temperature is concept in equilibrium state, so in non-equilibrium one, we had to assume local equilibrium as the assumption. In actual macroscopic climate system in years change would be so slow to enable the assumption. Note  $\sigma T_G(t)^4 \gg \delta F_0(t)$ .

$$C_G (dT_G(t)/dt) = 4\pi R^2 \delta F_0(t).$$

—heat budget diagram—



$$\sigma T_A(t)^4 \equiv \delta F_0(t) + \sigma T_G(t)^4. \rightarrow \delta F_0(t) = \sigma (T_A^4 - T_G(t)^4) \dots \dots \dots (6)$$

In the below, we assume step function change of GHG concentration from initial time  $t_0$  to final time  $t_E$ . In this interval, G could radiate and sink  $\sigma T_G(t)^4$ .

Also blackbody  $T_G(t)$  would be changed by heat input(output)  $\delta F_0(t)$  due to GHG concentration change, while  $T_A$  must be constant due to nothing heat reserve (output=input) during the interval. But it must that  $T_A = T_G(t_0)$  and  $T_A = T_G(t_E)$  due to equilibrium-ness between G & A at time  $t_0$  and  $t_E$ . It's very curious!. It's not a simple black body, but be a **pseudo** one of (6).

A with GHG has two kind of temperature, one is ordinal thermal one observed by thermometer, the other is radiative temperature represented by (6). Atmosphere radiation ( $\sigma T_A(t)^4$ ) are that of own molecular transit energy ( $\sigma T_G(t)^4$ ) with adding solar ray penetration of ( $\delta F_0(t)$ ).

## [2]:Estimation of the EGT solution:

(1)physical constants around EGT.

EGT is non-linear, but its derivative directly be calculated from right side term.

At first, we could prove being of **an inertia temperature rise with certain delay, even though we have fixwed past GHG concentration**. Therefore we need **more** (80%) **carbon reduction rate than that of {sink=emit} balanced amount of about 60%**.

$$\langle C_G \rangle = 1.4 \times 10^{24} \text{ J/k (ocean depth=1000m)}.$$

$R_E = 6.38 \times 10^6 \text{ m}$ =Earth radius,

$$\sigma_0 = \langle \text{Stefan Boltzmann} \rangle = 5.67 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4}.$$

$$* dT_G(t)/dt = (4\pi R_E^2 \sigma_0 / C_G) (T_A^4 - T_G(t)^4) \equiv K_G (T_A^4 - T_G(t)^4).$$

$$K_{G/t} \equiv (4\pi R_E^2 \sigma_0 / C_G) = 2.9 \times 10^6 / C_G (d) = 2.2 \times 10^{-17} (\text{d}=1000\text{m}),$$

(a)The ocean heat capacity  **$C_o = 5.3 \times 10^{24} \text{ J/K}$**  by linear estimation.

<http://hypertextbook.com/facts/2001/SyedQadri.shtml>. ocean volume= $1.37 \times 10^9 \text{ km}^3$ .

ocean weight of density  $1 = 1.4 \times 10^9 \text{ kg} (10^5 \text{ cm})^3 = 1.4 \times 10^{24} \text{ g}$ . specific heat of sea water = $3.85$  (pure water $4.18$ )  $\text{J/gK}$ .  $C_o = 3.85 \text{ J/gK} \times 1.4 \times 10^{24} \text{ g} = 5.3 \times 10^{24} \text{ J/K}$ .

(b)average depth of ocean= $3796\text{m}$ , solar ray reaching depth= $200\text{m}$ , then

roughly estimated active heat capa of ocean  $\approx 5.3 \times 10^{24} \text{ J/K} \times 200 / 4000 \approx 2.6 \times 10^{23} \text{ J/K}$ .

(c)Annual Effective global heat capacity  $C_G$  derived by NASA.

<http://www.ecd.bnl.gov/steve/pubs/HeatCapacity.pdf>.

Effective global heat capacity.  $C_e = 10^9 \text{ J m}^{-2} \text{ K}^{-1} \times 0.53 \pm 0.22$

$$C_G = 4\pi R_E^2 \times 0.53 \times 10^9 \text{ J m}^{-2} \text{ K}^{-1} = 4\pi (6.38 \times 10^6 \text{ m})^2 \times 0.53 \times 10^9 \text{ J m}^{-2} \text{ K}^{-1} = 2.7 \times 10^{23} \text{ J/K}.$$

(d)Long term effective global heat capaity  $C_G$  derived by radiative forcing.

$$* \delta F_0(t) = \sigma (T_A(t)^4 - T_G(t)^4).$$

$$* (4\pi R_E^2 \sigma_0 / C_G) \equiv K_G = [dT_G(t)/dt] / (T_A^4 - T_G(t)^4) = \sigma [dT_G(t)/dt] / \delta F_0(t).$$

$$\rightarrow C_G = 4\pi R_E^2 (\sigma_0 / \sigma) \langle \delta F_0(t) \rangle / \langle dT_G(t)/dt \rangle \approx 4\pi R_E^2 \delta F_0(t) / [dT_G(t)/dt].$$

$$C_G = 4\pi (6.38 \times 10^6 \text{ m})^2 \times 1.6 \text{ W m}^{-2} / [0.02 \text{ K} / 3600 \times 24 \times 365] = 1.29 \times 10^{24} \text{ J/K}.$$

It's about  $1000\text{m}$  depth by observed data of  $\{\delta F_0 \text{ & } dT_G/dt\}$ .

(e)time to year/variable transform: 1 year= $3600 \times 24 \times 365 \text{ sec} = 3.15 \times 10^7 \text{ s}$ .

$$K_{G/y} \equiv K_{G/t} \times 3600 \times 24 \times 365 = 7.09 \times 10^{-10}.$$

(2) : Final temperature rise  $T_A$  with fixed GHG concentration of past time  $t_0$  the estimation by observed current temperature = **287.5K** with the time derivative =  **$0.03 \sim 0.05\text{K/y}$** . <thermodynamical temperature  $273\text{K} \equiv 0^\circ\text{C}$ >

\*  $dT_G(t)/dt = (4\pi R_E^2 \sigma_0 / C_G) (T_A^4 - T_G(t)^4) \equiv K_G (T_A(m(t), b(t))^4 - T_G(t)^4)$  ... EGT.  
 $K_G \equiv (4\pi R_E^2 \sigma_0 / C_G) = 7.09 \times 10^{-10}$ : < $C_G$ =global heat capacity=oceans of depth 200m>

\*  $\sigma_0 T_A(m(t), b(t))^4 = F_0(1-m(t))/4(1-b(t))$ . <pseudo-final equilibrium temperature>  
**b**=heat flux trapp probability of GHG、**m**=albedo>

(a) a virtual fitting solution value with fixed  $T_A(m, b) = \text{constant}$ :

$$T(1750) = (273 + 13.7) = 286.7 ; T(2008) = (273 + 14.5) = 287.5 ;$$

288.7	286.7	286.9	287.1.	287.3	287.5
$dT/dt$	.074	.061	0.047	.034	0.020

☞: Caution that the actual trend( $dT_G/dt$ ) is obviously increasing as time goes on. That is, it is due to more increasing of  $T_A(m(t), b(t))$  due to "b" of GHG rise and albedo "m" decreasing. It means a going more away of "our relieve".

(b) Inertia temperature rise .

$$T_A = \{[(dT_G(t)/dt)/K_G] + T_G(t)^4\}^{1/4}.$$

$$\text{The final temperature} = \{\text{current trend}/K_G + (\text{now temperature})^4\}^{1/4}.$$

$$\text{current value : } T_G(t) = 277.5. \quad K_G = 4\pi R_E^2 \sigma_0 / C_G = 7.09 \times 10^{-10}.$$

$dT_G/dt$	$K_G$	$K_G(d=1000\text{m}) = 7.09 \times 10^{-10}$ :	
0.0177(IPCC)		recent 25 years trend, but now value is far from it !	
0.02K/y		$T_A = 277.83(+0.33\text{K})$	
0.03		277.99(+0.5K)	
0.04		278.16(+0.66K)	
0.05		278.32(+0.82K)	

(3) **Process time estimation:**  $\langle K_G = 4\pi R_E^2 \sigma_0 / C_G = 7.09 \times 10^{-10} \rangle$

(a) given  $T_A(t)$  the virtual final temperature as function of  $\{m(t), b(t)\}$ .

$$dT_G(t)/dt = K_G(T_A(t)^4 - T_G(t)^4). \rightarrow dt/dT_G = 1/K_G(T_A(t)^4 - T_G(t)^4).$$

(b) fixed  $T_A(t_0) > T_G(t) > T_G(t_0)$ .  $t - t_0 = \int_{T_A(t_0)}^{T_G(t)} dT_G / K_G(T_A(t_0)^4 - T_G^4)$ .

integral formula

$$1/(Q^4 - X^4) = (1/2Q^4) [1/(1 + (X/Q)^2)] + (1/4Q^3) \langle 1/(X+Q) - 1/(X-Q) \rangle.$$

$$\int dX/(Q^4 - X^4) = (1/2Q^3) \operatorname{Tan-1}(X/Q) + (1/4Q^3) \ln(|X+Q|/|X-Q|).$$

$$t - t_0 = \int_{T_G(t_0)}^{T_G(t)} dT_G / K_G(T_A(t_0)^4 - T_G^4);$$

$$T_G(t)/T_A(t_0), T_G(t_0)/T_A(t_0), |T_G(t) + T_A(t_0)|/|T_G(t_0) + T_A(t_0)| \approx 1 \text{ at near } 287.5\text{K}.$$

$$t - t_0 = (1/4K_G T_A(t_0)^3) \{ 2 \langle \operatorname{Tan-1}(T_G(t)/T_A(t_0)) - \operatorname{Tan-1}(T_G(t_0)/T_A(t_0)) \rangle + \ln(|T_G(t) + T_A(t_0)|/|T_G(t) - T_A(t_0)|) - \ln(|T_G(t_0) + T_A(t_0)|/|T_G(t_0) - T_A(t_0)|) \}$$

$$(1/4K_G T_A(t_0)^3) = 15y. \rightarrow t - t_0 \approx 15. \ln(|T_G(t_0) - T_A(t_0)|/|T_G(t) - T_A(t_0)|) \equiv 15W.$$

$$T_G(t_0) - T_A(t_0) \equiv 1.$$

$ T_G(t) - T_A(t_0) $	$W = \ln(1/ T_G(t) - T_A(t_0) )$	$(t - t_0)y$
0.9		1.6
0.8		3.4
0.5		10.4
0.2		24.32

$$T_G(t_0) - T_A(t_0) \equiv 2.$$

$ T_G(t) - T_A(t_0) $	$W = \ln(2/ T_G(t) - T_A(t_0) )$	$(t - t_0)y$
1.9		0.75
1.8		1.6
1.5		4.3
1.0		10.4
1.5		20.8

$$T_G(t_0) - T_A(t_0) \equiv 3.$$

$ T_G(t) - T_A(t_0) $	$W = \ln<3/ T_G(t) - T_A(t_0) >$	$(t - t_0)y$
2.9		0.5
2.8		1.1
2.5		2.8
2.0		6.1
1.0		16.5

While 0.5°C rise took (10, 5, 2.5) years,  $T_G$  is heading to (1, 2, 3)°C rise.

#### (4) The approximated solution $T_G(t)$ with $T_A(t)$ as the ZERO EMISSION TEMPERATURE.

Even though the ZERO CARBON EMISSION of most sink rate=1.5ppm/y seems risky.

(a) the global carbon budget:

[http://www.globalcarbonproject.org/carbonbudget/07/files/GCP\\_CarbonBudget\\_2007.Pdf](http://www.globalcarbonproject.org/carbonbudget/07/files/GCP_CarbonBudget_2007.Pdf)

fiscal (man made+natural) emit and (oceans+land) sink by photosynthesis.

+ man made emission = 7.5(8.5)PgC/y       $P=10^{15}$ . C is carbon standard,

+ natural emission = 1.5PgC/y.      For example)

- Oceans sinks = 2.3PgC/y       $CH_4=16g$ , but C=12g.

- Land sinks = 2.6PgC/y       $CO_2=44g$ , but C=12g.

+ atmospheric accumulation = 4.2PgC/y

current concentration rise : 4.2GtC=1.9ppm/y.

current manmade emit : 7.5GtC=3.6ppm/y.

net natural sink : 4.9GtC-1.5GtC=3.4GtC=1.5ppm/y. (MAX value of ZERO EMIT)

(b) current leading temperature  $T_A=287.8(+0.3K)$  with 385ppm CO<sub>2</sub> by  $dT_G(t)/dt=$

0.02~0.03K/y.

(c) coarse linear estimation on  $T_A(385\text{ppm}-1.5Y)$  the policy target at each year Y.

For the simplicity, leading temperature  $T_A$  is assumed as linear function of ppm.

Then 286.7(1750)K and 280ppm(1750) are equilibrium temperature and the

concentration in 1750 (Prior industrial revolution).

$$\begin{aligned} T_A(385\text{ppm} - (1.5\text{ppm}/\text{y})Y) &= 287.8 - (287.8 - 286.7) \times (1.5\text{ppm}/\text{y}) / (385\text{ppm} - 280\text{ppm}) \\ &= 287.8 - 1.1 \times 1.5Y / 105 = 287.8 - 0.016Y. \end{aligned}$$

(5)discrete integral of coarse approximantion on  $T_G(t)$  with the ZERO EMISSION.

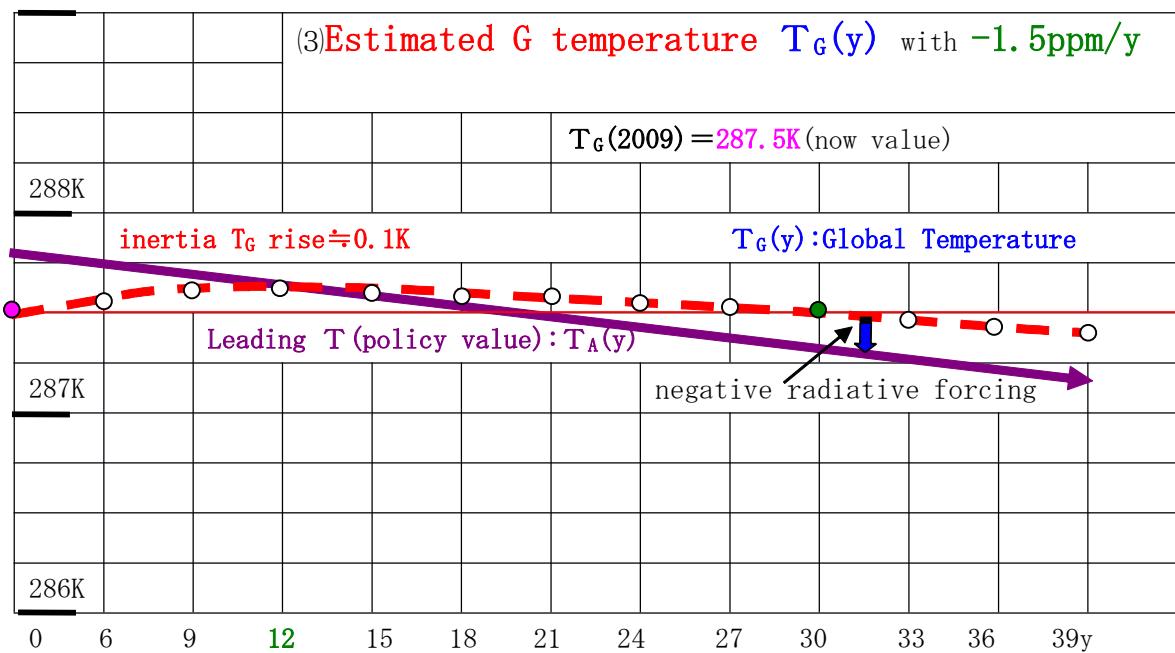
$$T_G(t + \Delta t) = T_G(t) + \Delta t (dT_G(t)/dt) = T_G(t) + \Delta t \cdot K_G(T_A(t)^4 - T_G(t)^4).$$

$$\Delta t = 3.0y (\Delta C = 4.5\text{ppm}) ; \Delta T = -0.016 \times 3 = -0.048\text{K}.$$

	385ppm	380.5	376	371.5	367	362.5	358	353.5
$\Delta y=3$	0y	3	6	9	12	15	18	21
$T_A$	<b>287.8</b>	287.75	287.70	287.66	287.61	287.56	287.51	287.46
$dT/dy$	<b>0.02K/y</b>	0.0128	<b>0.0067</b>	<b>0.0027</b>	<b>-.0013</b>	<b>-.0047</b>	<b>-.0067</b>	<b>-.0088</b>
$T' = dT+T$	<b><math>T_G=287.5</math></b>	287.56	287.60	287.61	287.63	287.62	287.61	287.59

	349	344.5	340	335.5	331	326.5	322	317.5
$\Delta y=3$	24	27	<b>30</b>	33	36	39	42	45
$T_A$	287.42	287.37	287.32	287.27	287.22	287.18	287.13	287.08
$dT/dy$	<b>-.0094</b>	<b>-.0101</b>	<b>-.0121</b>	<b>-.0128</b>	<b>-.0135</b>	<b>-.0134</b>	<b>-.0141</b>	<b>-.0148</b>
$T' = dT+T$	287.56	287.53	<b>287.50</b>	287.46	287.42	287.38	287.34	287.30

289K



(6)  $\{dT/dy=0.02K/y, T_A=287.8\}$  is conservative, let's try  $\{dT/dy=0.03K/y, T_A=288\}$ .

$$T_A(385\text{ppm} - (1.5\text{ppm}/y)Y) = 288.0 - Y(288.0 - 286.7) \times (1.5\text{ppm}/y) / <385\text{ppm} - 280\text{ppm}>.$$

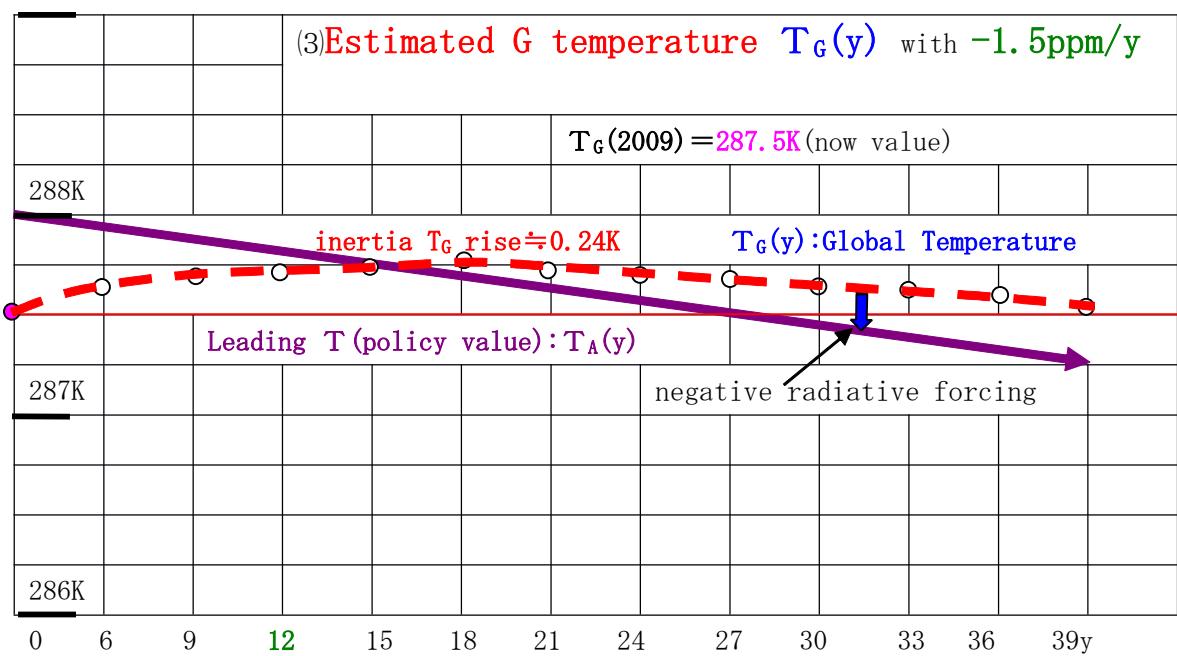
$$= 288.0 - 1.3 \times 1.5Y / 105 = 288.0 - 0.0186Y.$$

$$\Delta t = 3.0y, (\Delta C = 4.5\text{ppm}) ; \Delta T = -0.0186 \times 3 = -0.056K.$$

	385ppm	380.5	376	371.5	367	362.5	358	353.5
$\Delta y=3$	0y	3	6	9	12	15	18	21
$T_A$	288.0	287.94	287.89	287.83	287.78	287.72	287.66	287.61
$dT/dy$	0.03K/y	0.024	0.0155	0.0081	0.0034	-.0014	-.0054	-.0074
$T' = dT+T$	$T_G=287.5$	287.59	287.66	287.71	287.73	287.74	287.74	287.72

	349	344.5	340	335.5	331	326.5	322	317.5
$\Delta y=3$	24	27	30	33	36	39	42	45
$T_A$	287.55	287.50	287.44	287.38	287.33	287.27	287.22	287.16
$dT/dy$	-.010	-.011	-.013	-.015	-.015	-.017	-.017	-.017
$T' = dT+T$	287.70	287.67	287.64	287.60	287.56	287.52	287.47	287.42

289K



### [ 3 ]:supplemental comment:

(1)The actual  $T_G(t)$  is almost exponential increasing which means that the leading temperature  $\equiv T_A(m(t), b(t))$  is becoming more far from the current  $T_G(t)$ . If the estimated leading temperature  $T_A$  be aheading  $0.5^\circ\text{C}$ , it be too terribly risky. Then a possible relief may be the maximum carbon sink rate=1.5ppm/y(ZERO EMIT). Even though, there would be possible natural spontaneous emit during temperature rise, so the final relieve may be nothing without global scaling forestization with rapid growing vegetations and pray God. Global climate-engineering ?, the action and the reaction ?.

(2)According to a coarse approximation of segmented integral calculation on EGT, it could scarcely recover current temperature 287.5K by taking 30 years with the maximum canbon sink rate 1.5ppm/y(the CARBON ZERO EMISSION). Then the max inertia temperature rise is  $\approx 0.1^0.3\text{K}$ , while we could only pray for nothing unfortunate. If possible safety temperature rise margin(PSTRM) be  $0.1^0.3^\circ\text{C}$  for preventing starting full dress positive feedback of natural spontaneous GHG emission, the current state is terribly risky !!!.

(3)GHG concentrations= $\{C_k(t)\}$  and super black body factor $\equiv b(t)$ :<revised>.

<http://www.geocities.jp/sqkh5981g/RADIATION-FORCE.pdf>

The result mentioned in above page was revised!. Current radiative forcing  $\delta F_0(t)$  is fuction of  $b(t)$  which is also function of each GHG components concentrations  $\{C_k(t)\} \rightarrow b(t)=b(C_k(t))$ . This is the problem to be solved in general.

$$\frac{1}{4}F_0(1-m(t))/(1-b(t)) = \sigma_0 T_A^4(m(t), b(t)) \equiv F(t).$$

$$C_G(dT_G(t)/dt) = 4\pi R^2 \delta F_0(t) = 4\pi R^2 \sigma_0 (T_A(t)^4 - T_G(t)^4).$$

☞:The definition "radiative forcing of climate betweenn 1750 and 2005 ( $1.6\text{W/m}^2$ )" by IPCC is not correct. But "betweenn 1750 and 2005"→"at now" is correct.

$(=1.6\text{W/m}^2)$  could be observed by between satelite and the ground.

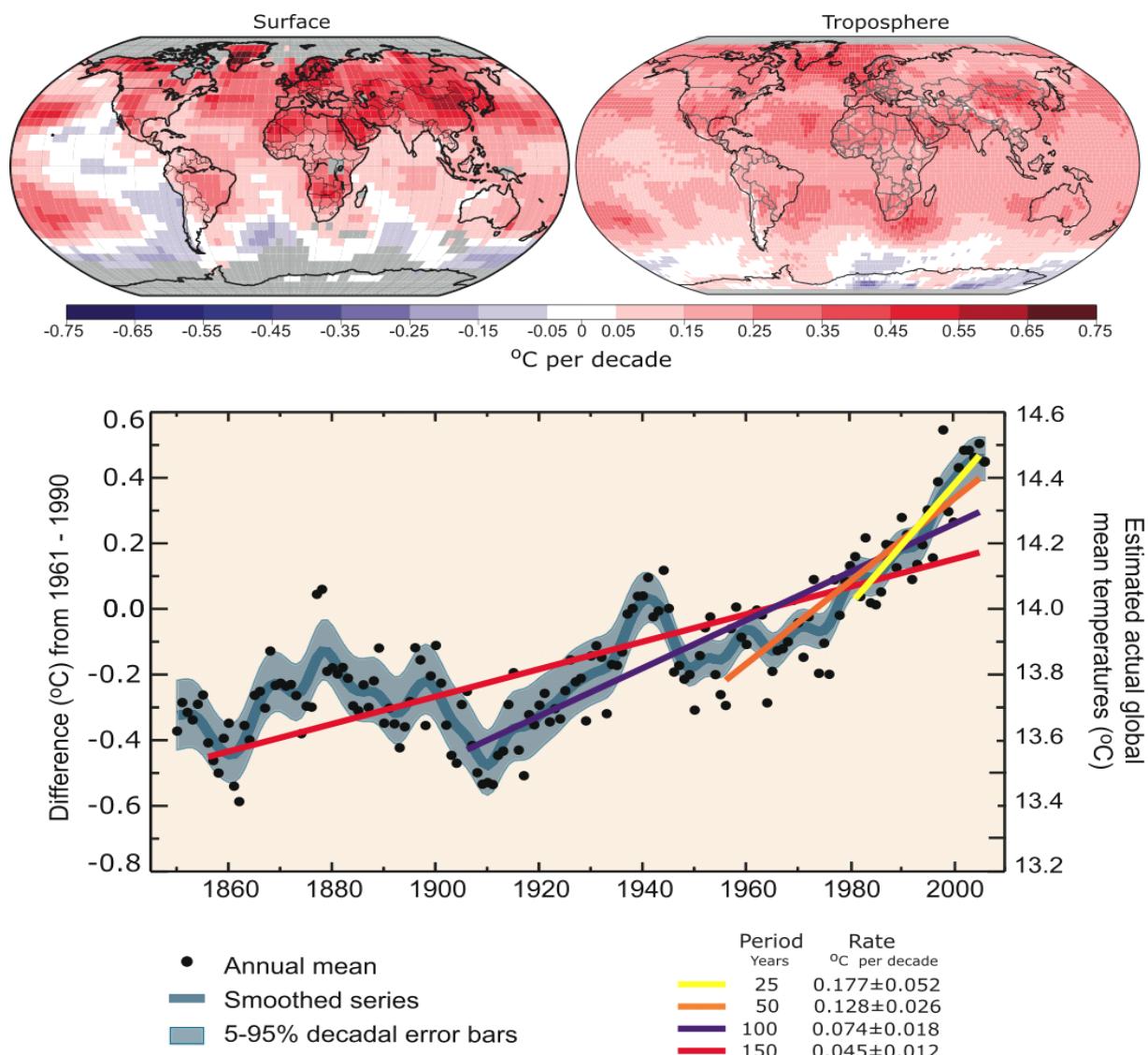
<http://www.ipcc.ch/pdf/assessment-report/ar4/wg1/ar4-wg1-chapter2.pdf>(p136).

If actual ( $\delta F_0=1.6\text{W/m}^2$ ) were different, the situation became serious.

**Appendix1 : Global Temperature & Carbon concentration Record by IPCC.**

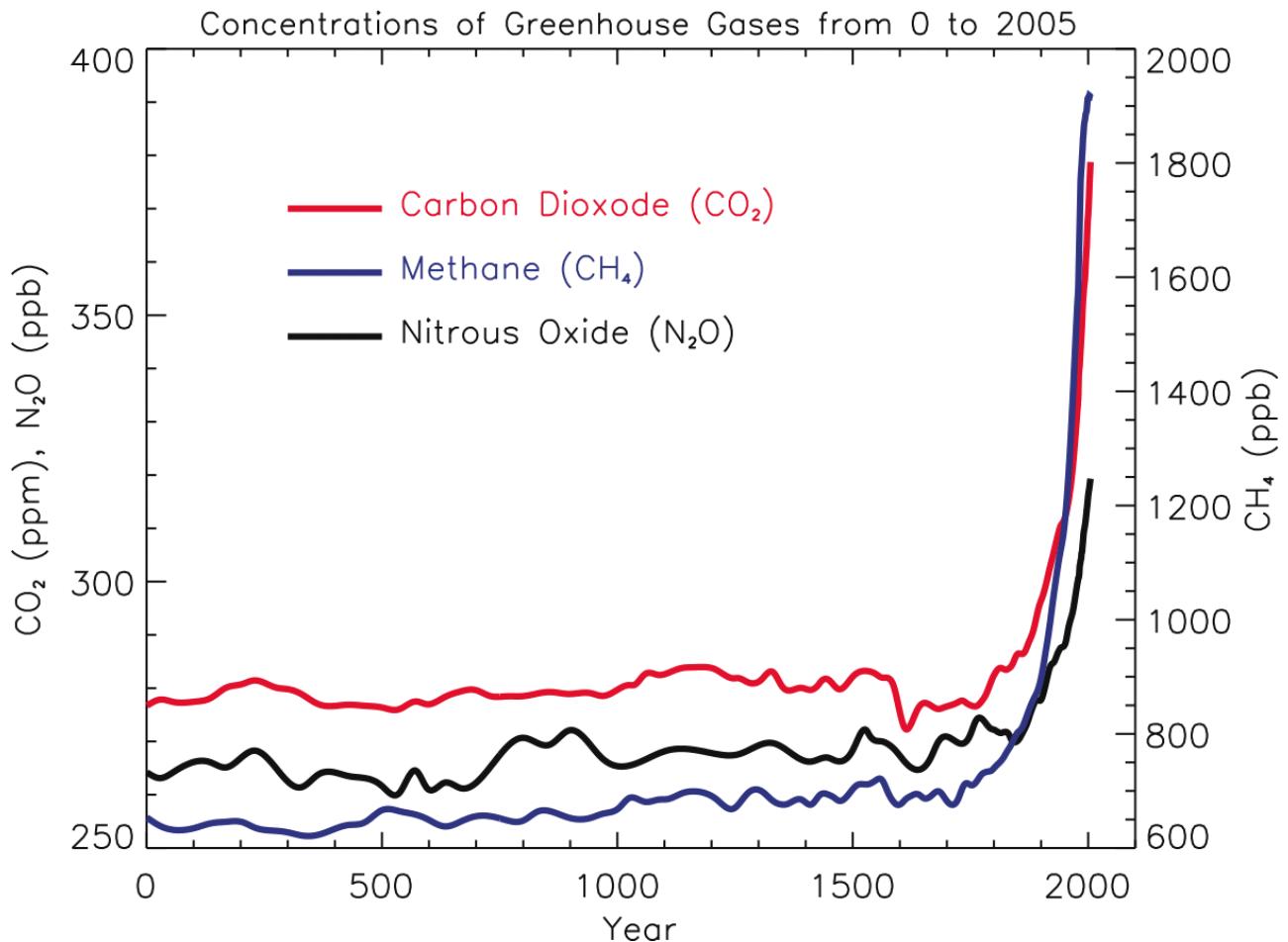
**IPCC Working Group I;Technical Summary;FINAL FIGURES**

**Historical Record of Global Temperature (p7)**

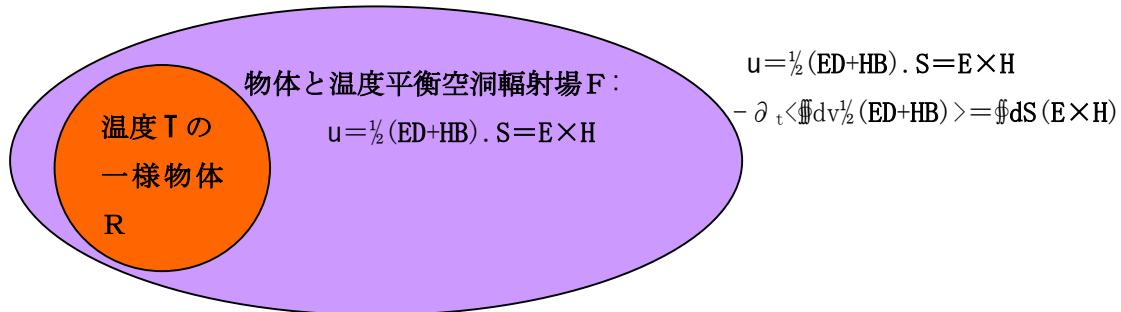


IPCC Working Group I;Technical Summary;FINAL FIGURES

Historical Record of Global Carbon concentration (p26)



## 付録2 : Planck 分布と非黒体 Stefan-Boltzmann 定数: (rev' 09/10/7)



(1) R 物性固有の温度 T 輻射密度分布 :  $R(E; T) = R(E = h\nu; T), 1 = \int_0^\infty dE R(E = h\nu; T).$

(2) energy 密度 :  $\rho = \int_0^\infty dE \cdot E \cdot R(E = h\nu; T) \equiv \sigma_R T^4.$

(3) Planck 分布 :  $R_0(E; T) = D(E) / (\exp(\beta E) - 1) = (8\pi E^2/c^3 h^3) / (\exp(\beta E) - 1).$

(a) 電磁場縮退度 D(E) の算出 :  $\langle L = 2r \text{ 球体空洞での } \lambda = c/\nu \text{ の電磁場姿態数を勘定} \rangle.$

(b) 立方体  $L = (\lambda_j/2)n_j, \langle n_j = 1, 2, 3, \dots \rangle (j=1, 2, 3); \nu_j = c/\lambda_j.$

$$\nu = (c/2L) \sqrt[n]{[n_j^2 + n_j^2 + n_j^2]} = (c/2L)n_j; n_j^2 + n_j^2 + n_j^2 = n^2.$$

(c) 振動数  $\nu \sim \nu + d\nu$  にある姿態度数  $= D(\nu) d\nu$  を勘定する。

「半径  $n$  の球体の  $1/8$  の内部の格子点と進行後退波 2 個を勘定。」

格子点は単位体積に付き、一個の勘定になる。だから」

$$N_n(\nu) = (1/4)(4/3)\pi n^3 = (1/4)(4/3)\pi (2L\nu/c)^3 = (8/3)\pi (L\nu/c)^3.$$

$$dN_n/d\nu = 8\pi (L/c)^3 \nu^2. \rightarrow D(\nu) = (dN_n/d\nu)/L^3 = 8\pi \nu^2/c^3.$$

(d) 完全黒体輻射分布 :  $R(\nu) d\nu = d\nu \cdot 8\pi \nu^2/c^3 (\exp(h\nu/kT) - 1).$

$$R(E) dE = dE \cdot 8\pi E^2/h^3 c^3 (\exp(\beta E) - 1).$$

$$(e) \rho = \int_0^\infty dE R(E) = \int_0^\infty dE \cdot 8\pi (\beta E)^3/h^3 (\beta c)^3 (\exp(\beta E) - 1)$$

$$= \langle 8\pi/h^3 \beta^4 c^3 \rangle [ \int_0^\infty d(\beta E) \cdot (\beta E)^3 / (\exp(\beta E) - 1) ] = [\pi^4/15] \langle 8\pi/h^3 \beta^4 c^3 \rangle$$

(e) 完全黒体物質と温度平衡にある輻射場 energy 密度 :

$$\rho = [8\pi^5/15] T^4 k_B^4/h^3 c^3 = (7.56 \times 10^{-16} \text{ J m}^{-3} \text{ K}^{-4}) T^4 = \frac{1}{2} (ED + HB).$$

$$\sigma_0 = [8\pi^5/15] k_B^4/h^3 c^3 = 7.562 \times 10^{-16} \text{ J m}^{-3} \text{ K}^{-4}. \langle \text{Stefan-Boltzmann 定数} \rangle$$

(f) 輻射場 energy 流密度 :

$$(E \times H) = j = (1/4)c\rho \text{ (従来論)} = \sigma_0 T^4. \quad \langle (11/48)c\rho \text{ (筆者 2008)} \rangle.$$

<http://www.geocities.jp/sqkh5981g/cu11D48.pdf>

$$\sigma_0 = (c/4) \sigma_0 = [2\pi^5/15] k_B^4/h^3 c^2 = 5.67 \times 10^{-8} \text{ J m}^{-3} \text{ K}^{-4}. \langle \text{大勢の Stefan-Boltzmann 定数} \rangle$$

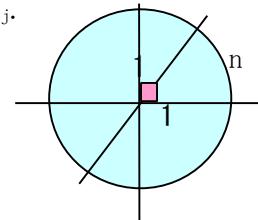
MKSA 単位:  $c = 3.00 \times 10^8 \text{ ms}^{-1}; h = 6.63 \times 10^{-34} \text{ Js}; k_B = 1.38 \times 10^{-23} \text{ J deg}^{-1}; F \text{ (太陽定数)} = 1368 \text{ W/m}^2.$

(g) 非黒体輻射 : 空中等の SB 定数(真空中とほぼ同じ) :

$$c = 2.99792459 \times 10^8 \text{ ms}^{-1}; = 1/\sqrt{\epsilon_0 \mu_0}. \quad c_a = 2.99704764 \times 10^8 \text{ ms}^{-1}; = 1/\sqrt{\epsilon_0 \mu_0}.$$

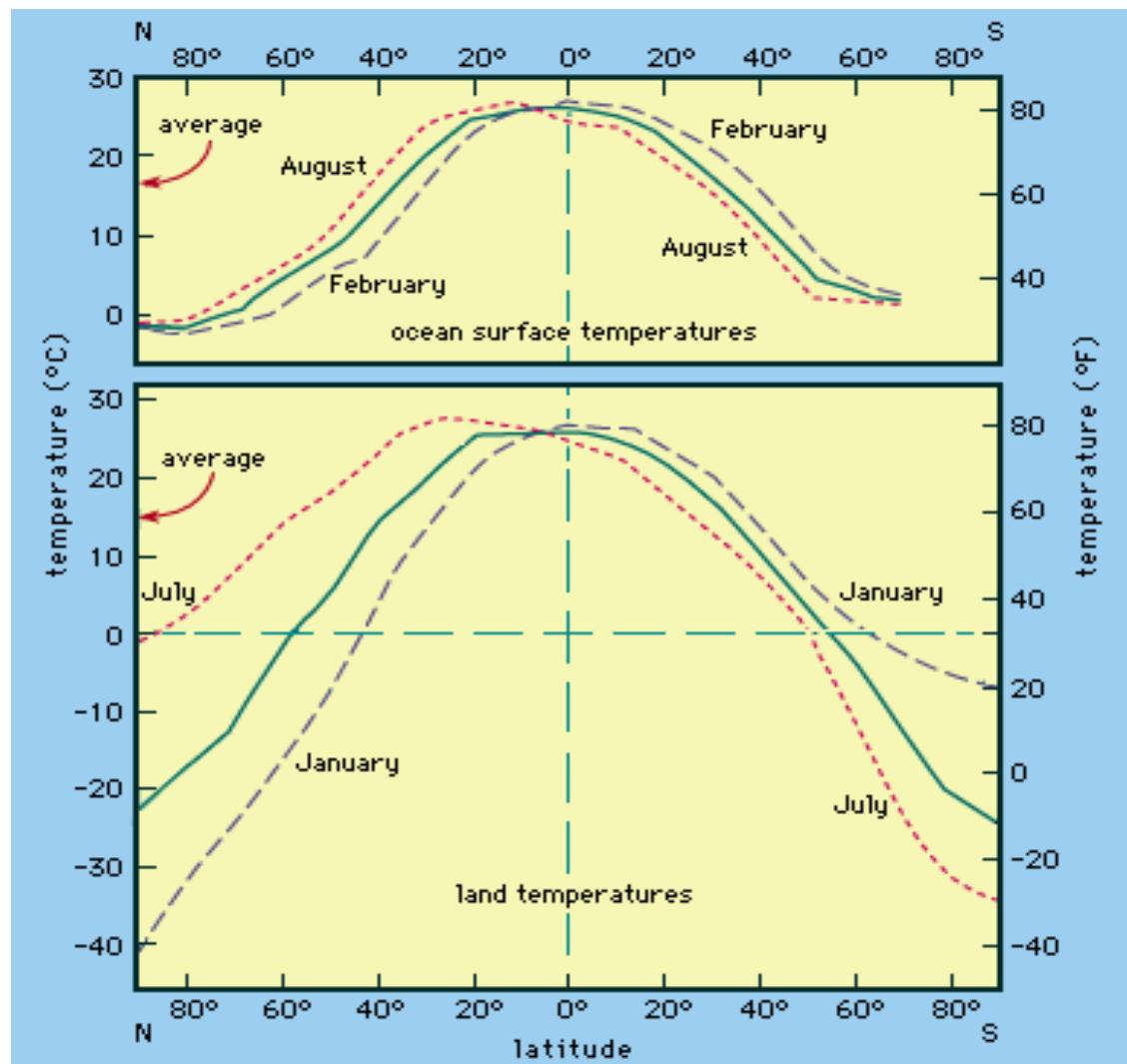
$(c_a/c)^3 = 1.00087$ , SB 定数は水蒸気があると波長短縮で大きくなり得る !!

\* 水中の光速度(比誘電率 80.4)。



### Appendix 3 :The Current Averaged Globala Ocean Temperature.

<http://www.britannica.com/EBchecked/topic/424285/ocean/67078/Temperature-distribution>



-postscript (ver5:' 09/11/20)-

This is also a cut and try paper in re-examination for radiative forcing.

In "version 5", the serious factors are corrected.

$C_G = 1.29 \times 10^{24} \text{ J/K}$  ?  $\leftarrow 2.7 \times 10^{23} \text{ J/K}$ .

$K_G = 7.09 \times 10^{-10}$  ?  $\leftarrow 3.5 \times 10^{-10}$ .

$dT_G/dy = 0.02 \sim 0.03 \text{ K/y}$  ?  $\leftarrow 0.05 \text{ K/y}$ .

and the conclusions such as prediction on global temperature are largely corrected. Before long(?), more-revised version would be presented.

e-mail: [motoji-suzuki@key.ocn.ne.Jp](mailto:motoji-suzuki@key.ocn.ne.jp)